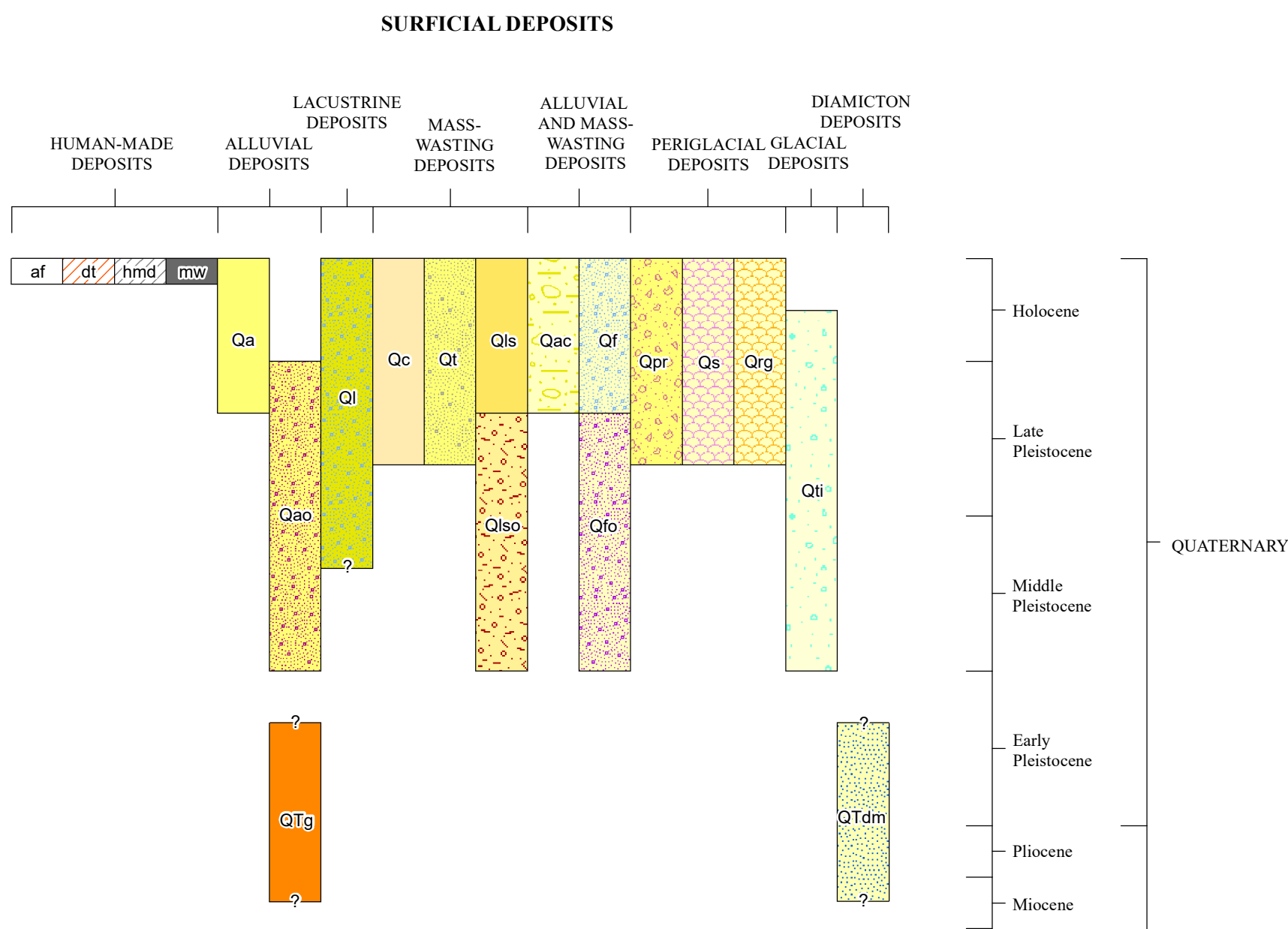
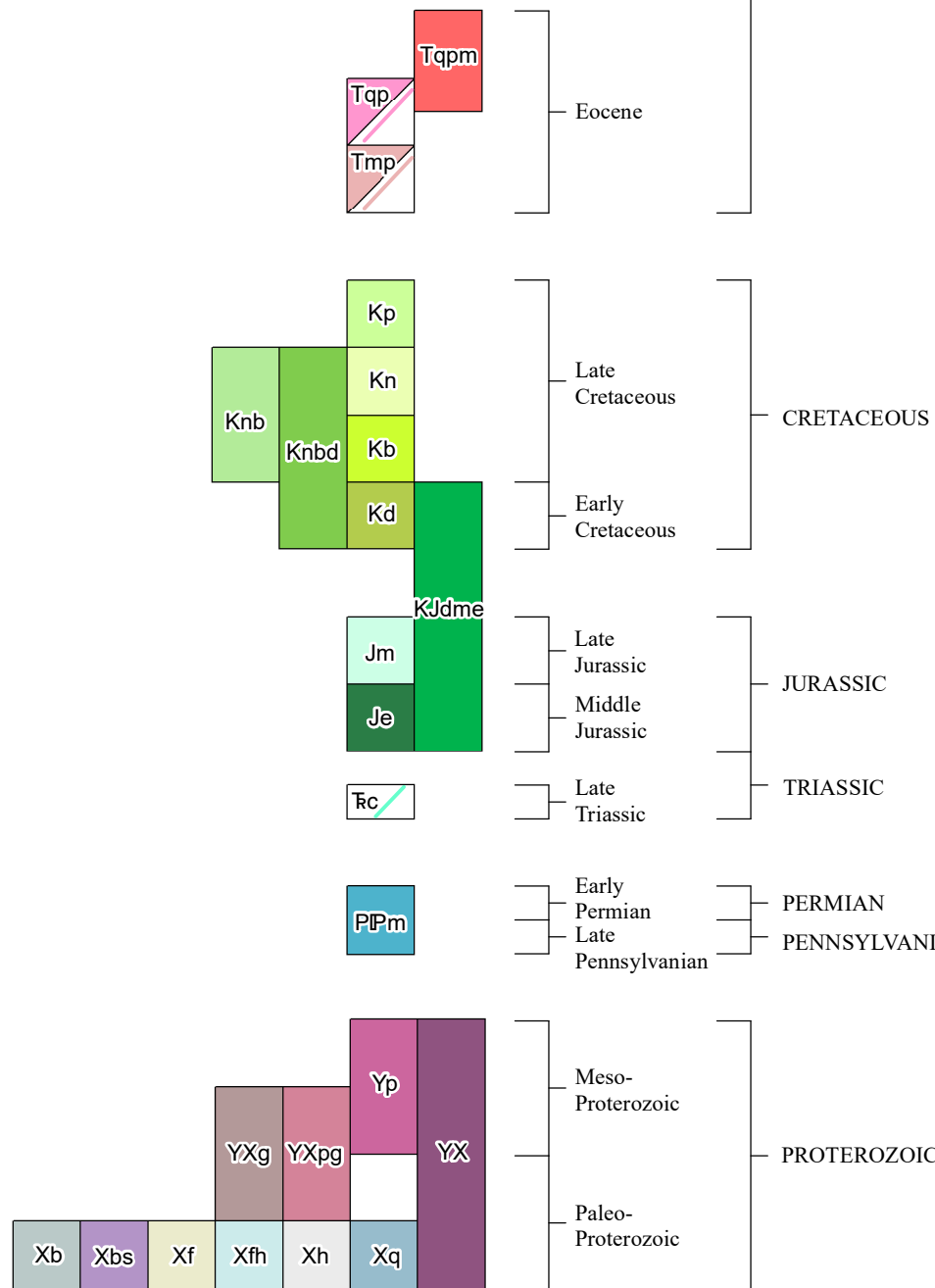


CORRELATION OF MAP UNITS



BEDROCK GEOLOGY



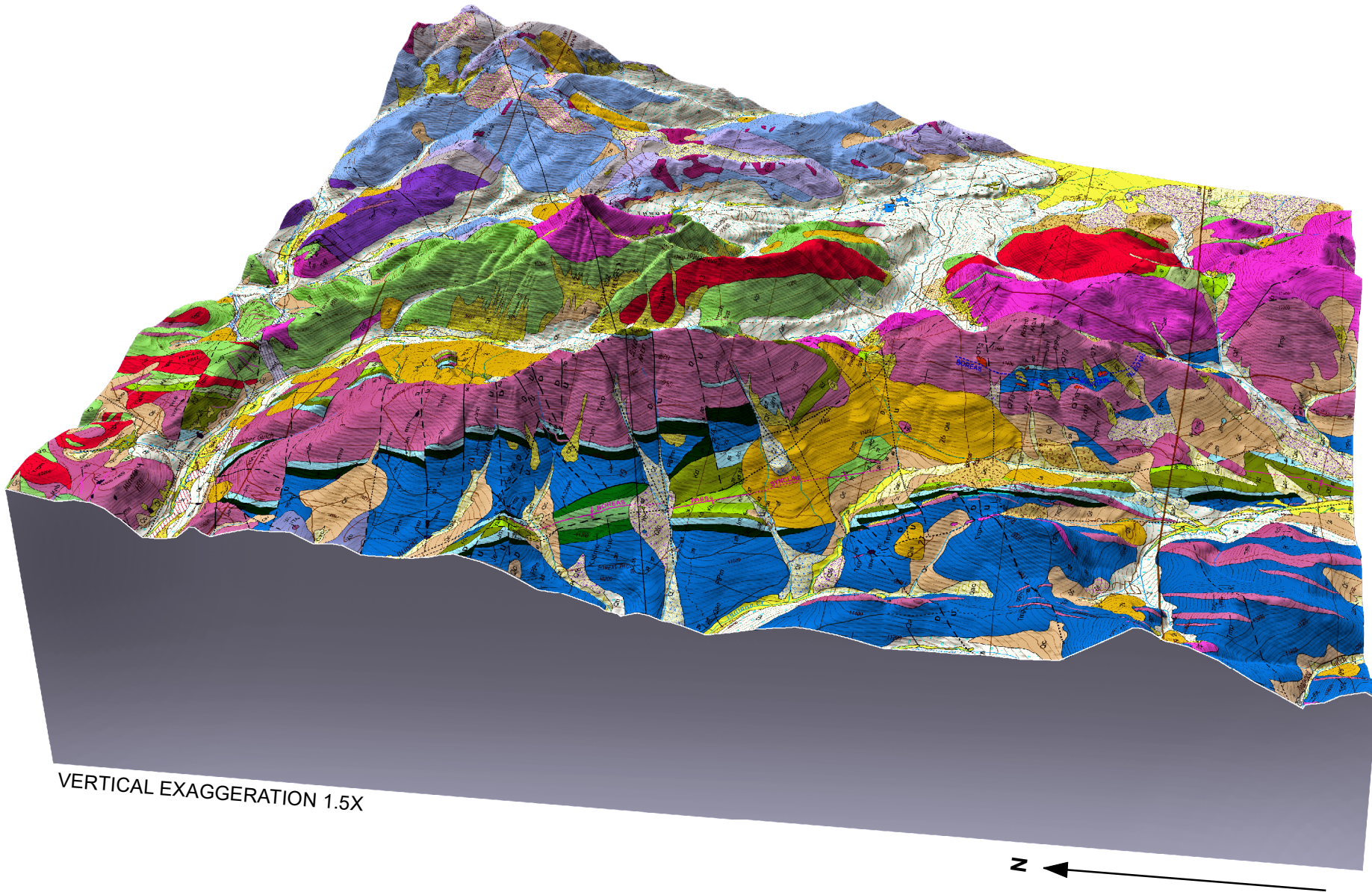
⁴⁰Ar/³⁹Ar AGE DATE INFORMATION

Results	400/(390) ± 3σ	403/(393) ± 3σ	Age ± 3σ	Mean	35Ar/(t) [%]	k/Ca ± 3σ
Age Plateau	15.4210 ± 0.0066	45.27 ± 0.12	1.41	50.18	13%	6.4 ± 0.5
	± 0.0066	± 0.20%				
	Ful External Error ± 1.02	± 0.20%	1.73	3σ Confidence Limit		
	Analyst Error ± 0.02	± 0.02	1.38%	Error Magnification		
Total Fusion Age	15.4029 ± 0.0038	46.32 ± 0.12	40	8.7 ± 0.0		
	± 0.0038	± 0.02				
	Ful External Error ± 1.05	± 0.02				
	Analyst Error ± 0.01	± 0.01				
Normal Isochron	234.2 ± 6.21	15.4629 ± 0.0126	46.32 ± 0.12	0.84	10.14	
	± 2.13%	± 0.008%	± 0.20%	10%	1%	
	Ful External Error ± 0.51	± 0.01	± 0.20%	3.28	3σ Confidence Limit	
	Analyst Error ± 0.04	± 0.04	1.000%	Error Magnification		
Inverse Isochron	286.40 ± 6.21	15.4589 ± 0.0026	46.31 ± 0.12	0.85	98.18	
	± 2.13%	± 0.008%	± 0.20%	10%	1%	
	Ful External Error ± 1.02	± 0.01	± 0.20%	3.28	3σ Confidence Limit	
	Analyst Error ± 0.06	± 0.06	1.000%	Error Magnification		

Results	400/(390) ± 3σ	403/(393) ± 3σ	Age ± 3σ	Mean	35Ar/(t) [%]	k/Ca ± 3σ
Age Plateau	15.3110 ± 0.0418	47.84 ± 0.18	17.14	55.70	7%	0.002 ± 0.002
	± 0.27%	± 0.18%				
	Ful External Error ± 1.09	± 0.18%	2.13	3σ Confidence Limit		
	Analyst Error ± 0.14	± 0.14	1.00%	Error Magnification		
Total Fusion Age	15.2705 ± 0.0090	47.48 ± 0.12	28	63.04 ± 0.00		
	± 0.06%	± 0.06%				
	Ful External Error ± 1.08	± 0.06%	2.13	3σ Confidence Limit		
	Analyst Error ± 0.07	± 0.07	1.00%	Error Magnification		
Normal Isochron	154.28 ± 239.62	15.3777 ± 0.0131	47.47 ± 0.08	20.00	55.70	
	± 41.23%	± 0.008%	± 0.08%	0%	2σ Confidence Limit	
	Ful External Error ± 1.51	± 0.01	± 0.08%	4.979	Error Magnification	
	Analyst Error ± 1.46	± 1.46				
Inverse Isochron	154.28 ± 122.07	15.3681 ± 0.0074	47.43 ± 0.07	20.31	55.70	
	± 39.80%	± 0.007%	± 0.07%	0%		
	Ful External Error ± 1.52	± 0.01	± 0.08%	4.979	Error Magnification	
	Analyst Error ± 1.47	± 1.47				

Figure 10 consists of two panels, (a) and (b), each showing a plateau of age data and associated isochron plots. Panel (a) shows a plateau at 45.27 ± 0.12 Ma. The top plot is a plateau of age data (Age [Ma] vs Cumulative 35Ar Released [%]) with a horizontal line at 45.27 ± 0.12 Ma. The bottom plot is a 35Ar/40Ar vs 39Ar/40Ar plot showing a linear decrease in 35Ar/40Ar with increasing 39Ar/40Ar. Panel (b) shows a plateau at 47.84 ± 0.18 Ma. The top plot is a plateau of age data (Age [Ma] vs Cumulative 35Ar Released [%]) with a horizontal line at 47.84 ± 0.18 Ma. The bottom plot is a 35Ar/40Ar vs 39Ar/40Ar plot showing a linear decrease in 35Ar/40Ar with increasing 39Ar/40Ar.

3-D OBLIQUE VIEW OF GEOLOGIC MAP



GEOLOGIC OVERVIEW

The Borcas Pass quadrangle, located entirely above elevation 9,800 ft (2987 m), is in the far southwestern part of the Front Range, a major constituent of the southern Rocky Mountains. A small area in the southeastern map area is in the northwestern part of South Park, a large, structurally complex intermontane basin. Geologically, the quadrangle is very complex, having experienced several periods of structural deformation and uplift, igneous intrusive activity, hydrothermal alteration and mineralization, and intense alpine glaciation.

Paleoproterozoic amphibolite-facies layered metamorphic rocks and Paleoproterozoic to Mesoproterozoic granitic intrusive rocks are exposed in the northeastern part of the quadrangle and in a small area along the western edge. The layered metamorphic gneisses and rare schists were derived from dominantly clastic sedimentary protoliths, possibly with some interlayered volcanic rocks. The layered rocks are foliated and commonly folded. Regional metamorphism and deformation occurred concurrently with the emplacement of syntectonic plutons during the Paleoproterozoic (Reed and others, 1987).

During the early to middle Paleozoic, the Proterozoic rocks were likely buried by marine clastic and carbonate sedimentary rocks. Early Paleozoic rocks of this type are present in the Ancestral Rocky Mountains (see, for example, the review by Williams, 1972) quadrangles (Wallace and others, 2002; Widmann and others, 2004). All of these early and middle Paleozoic marine sedimentary rocks were eroded away, at least in the northern and eastern map area, when the Ancestral Rocky Mountains were uplifted and eroded. The Ancestral Rocky Mountains formed the red bed sequences of the Pennsylvanian and Permian Maroon Formation found in the quadrangle. The Maroon Formation thickens rapidly from north to south across the quadrangle. It is less than 600 ft thick (183 m) in the northwestern part of the map area and adjacent Breckinridge quadrangle. The southwestern part of the map area and the adjacent Breckinridge quadrangle. The Maroon Formation may never have been deposited in the northeastern map area. This indicates that the quadrangle may have straddled the boundary between a highland area of the Ancestral Rocky Mountains and the deep Central Colorado Trough (De Yoo, 1972) sedimentary basin.

The Triassic Chinle Formation and the colian Middle Jurassic Entrada Sandstone were deposited unconformably over an irregular erosion surface that developed on the Maroon Formation. Further deposition of sedimentary rocks, separated at times by periods of non-deposition and/or erosion, characterized the Late Jurassic to Late Cretaceous periods here as elsewhere in Colorado. The Western Interior Seaway covered the area during Cretaceous time and deposited the clastic and carbonate rocks of the Dakota Sandstone, Niobrara Formation, Benton Shale, and Pierre Shale found in the quadrangle. Combined, the Cretaceous sedimentary rocks are estimated to be over 2,500 ft thick (762 m) in the map area.

All of the Proterozoic through Upper Cretaceous rocks in the region were completely deformed and uplifted during the Laramide orogeny, which began near the end of the Cretaceous Period. A major regional thrust fault, the Williams Range thrust, transects the quadrangle from northwest to southeast and is interpreted to connect with the Elkhorn thrust fault in South Park. The thrust fault system is fairly well exposed in the northern part of the map area; however, farther south it is completely covered by Quaternary surficial deposits for over four miles in the Michigan Creek drainage. The Williams Range thrust tectonically placed Proterozoic metamorphic and igneous rocks over the Cretaceous Pierre Shale. Many other faults in the quadrangle, as well as the two mapped synclines, also formed during the Laramide orogeny.

In middle Eocene rocks, which may correspond to the latter stages of the Laramide orogeny in central Colorado, porphyritic intrusions of intermediate to silicic composition were emplaced along and west of the Williams Range thrust fault zone. These intrusions include monzonite porphyry (Mtp) and quartz monzonite porphyry (Qmzp), both of which are less than 5 ft (1.5 m) to over 200 ft (60 m). Ne "Ar"/Ar age dating yielded an age of 47.84 ± 0.18 Ma for monzonite porphyry (Tmnp) and 45.27 ± 0.12 Ma for quartz monzonite porphyry (Tqmp). Uplift caused by the emplacement of intrusions may have caused some shallow dipping Williams Range thrust faults upward, especially near the Mt Grouse area. The fact that no faults were observed to the northeast, and that there was one unambiguously dated middle Eocene dikes were observed, and these were in Proterozoic rocks close to the Williams Range thrust. However, porphyry dikes (typically less than 20 ft (6 m) thick) were reported in mine workings in the historic Breckenridge mining district, the eastern portion of which is in the northwestern part of the study area. These dikes were probably related to the same magmatic system as the economically significant quantities of gold, silver, lead, and zinc in the Breckenridge district. These deposits are interpreted to be genetically related to the Eocene igneous intrusive activity. Deposits were focused along faults, fractures, and contacts that formed open spaces and locally enhanced the mineralizing hydrothermal fluids. Superheated steam-heated rocks derived from the continuous vents and vented gases at shallow, but locally spectacularly rich deposits near the surface that were worked using hydraulic and placer methods, including those in American, Georgia, Dry, and Monitor gulches, and at the Wire Patch deposit on the northwest slope of Farncomb Hill. These deposits are interpreted to be genetically related to the Eocene igneous intrusive contact metamorphic mineral deposits (skarns) are present in thin calc-silicate-altered carbonate beds in sedimentary rocks adjacent to or very near intrusives.

Numerous faults, mostly vertical displacements of less than 1000 ft (305 m), cut the Eocene intrusives as well as older rocks, especially in the northern half of the quadrangle. Most of these strike east-northeast to northeast. From late Tertiary (Neogene) time to the present, erosion of the high mountains was the dominant geological process in the Boreas Pass region. Evidence of a late Tertiary or early Pleistocene paleovalley was found on the high western and southern slopes of Boreas Mountain where small remnants of coarse, fluvial gravel are present. A large diamictite, possibly as old as Pliocene, composed of mostly sub-angular gravel was discovered along and south of the Continental Divide east of Georgia Pass. Further work is needed to determine the origin of this gravel, which locally forms the upper half of the Quaternary glacial cirque.

During Quaternary time, glaciers produced numerous cirques and the deeply glaciated valleys of French Gulch and the Middle and South Forks of the Swan River on the northwest side of the Continental Divide, and Michigan Creek, French Creek, Jefferson Creek, and the Tarryfork Creek on the east. Deposition of glacial till and outwash were produced. Rock glaciers are present in the South Fork of the Swan River at the high cirque basins. A few visually spectacular proglacial rampart deposits were formed from loose rocks falling and being deposited as arcuate ridges along the bases of long-lasting snow and ice fields. Rich placer gold deposits formed during the Pleistocene in the French Gulch and the Middle and South Forks of the Swan River. The map area of the map area. Placers in the non-glaciated, smaller gulches north and east of Farncomb Hill were formed by mass-wasting and erosion of supereigne deposits. Landslides cover a substantial portion of the Boreas Pass quadrangle and are the most common natural hazard. Other geologic hazards in the map area include rockfall, debris flows, floods, and swelling soils. No evidence of Quaternary faulting was recorded in the map area.

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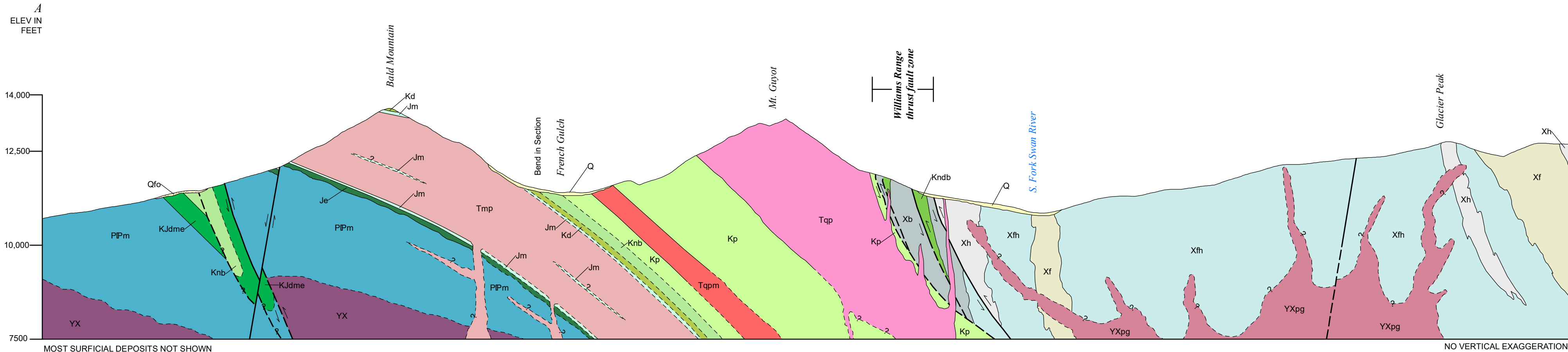
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CROSS SECTION A-A'



GEOLOGIC MAP OF THE BOREAS PASS QUADRANGLE, PARK AND SUMMIT COUNTIES, COLORADO
CORRELATION OF MAP UNITS, 3-D OBLIQUE MAP, GEOLOGIC OVERVIEW, AND CROSS SECTION

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