

CONDENSED DESCRIPTION OF MAP UNITS

The complete description of map units and references is in the accompanying booklet.

SURFICIAL DEPOSITS

HUMAN-MADE DEPOSITS

- af** Artificial fill (late Holocene)—Unsorted silt, sand, and rock fragments from alluvial, debris-fan, and rockfall deposits, mine and mill waste, and recycled construction materials deposited by humans
- mw** Mine and mill waste (late Holocene)—Pebble- to cobble-size waste rock from mines, prospecting pits, and milling operations

ALLUVIAL DEPOSITS—Sediments deposited in stream channels, on flood plains, terraces, and minor debris fans, and in sheet-wash areas

- Qa** Stream-channel, flood-plain, and low terrace alluvium (Holocene)—Mostly clast-supported, sandy to bouldery pebble and cobble gravel interbedded and often overlain by silty sand. Includes organic-rich mud, peat, sand, and gravel underlying terraces up to about 10 ft above modern stream level

- Qas** Older alluvium (late Pleistocene)—Poorly sorted, matrix-supported, pebble and cobble gravel in a sandy matrix. Includes stream alluvium and sheetwash deposits underlying a single small terrace remnant near Green Lake about 40 ft above the modern South Clear Creek stream level

COLLUVIAL DEPOSITS—Silt, sand, and gravel on valley sides and floors transported and deposited primarily by gravity

- Qc** Colluvium (Holocene and late Pleistocene)—Ranges from clast-supported, pebble to boulder gravel in a sandy matrix to matrix-supported gravely, silty sand
- Qls** Landslide deposits (Holocene and late Pleistocene)—Clast-supported, heterogeneous deposits consisting primarily of large blocks of rock and silt- to boulder-size debris. Includes slides characterized by numerous linear scarps, hummocky terrain, interrupted drainages, and large terrace blocks
- Qt** Talus deposits (Holocene and late Pleistocene)—Bedrock-derived angular, cobbly and bouldery rubble transported as rockfalls, rockslides, rock avalanches, and rock topples

ALLUVIAL AND COLLUVIAL DEPOSITS—Sediments in stream channels and on flood plains, debris fans, and lower reaches of adjacent hillslopes

- Qf** Debris-fan deposits (Holocene and late Pleistocene)—Poorly sorted to moderately sorted, matrix-supported, gravely, sandy silt to clast-supported, pebble and cobble gravel in a sandy silt or silty sand matrix. Includes debris-flow, hyperconcentrated-flow, fluvial, and sheetwash deposits on debris-fans and in drainage channels
- Qac** Alluvium and colluvium, undivided (Holocene and late Pleistocene)—Poorly sorted to moderately sorted, stratified, interbedded sand, pebbly sand, and sandy gravel to poorly sorted, unstratified or poorly stratified, silty sand and bouldery sand
- Qds** Colluvium and sheetwash deposits, undivided (Holocene and late Pleistocene)—Colluvium on steeper slopes, sheetwash on gentler slopes. Colluvium ranges from unsorted to poorly sorted, clast-supported, pebble to boulder gravel in a sandy matrix to matrix-supported gravely sand. Sheetwash deposits typically consist of pebbly, silty sand and sandy silt

PERIGLACIAL DEPOSITS—Boulder- to clay-size fragments transported primarily by solifluction and creep

- Qe** Solifluction deposits (Holocene and late Pleistocene)—Pebbles, cobbles, and large boulders in a sandy matrix transported down slope by solifluction and frost creep

GLACIAL DEPOSITS—Gravel, sand, silt, and clay deposited by

- Qe** Till (late and late middle Pleistocene)—Unsorted, unstratified, matrix-supported bouldery to pebbly gravel in a sandy matrix deposited by ice. Locally may contain interlayers of stratified drift

NONSORTED TERRIGENOUS DEPOSITS—Bouldery material of uncertain origin

- Qe** Diamicton (early to middle Pleistocene)—Non-sorted deposits consisting largely of boulders and cobbles in a sandy matrix. May have been deposited by pre-Bull Lake ice, landslides, earthflows, mudflows, or solifluction (Plint and others, 1960)

BEDROCK

TERTIARY AND UPPER CRETACEOUS INTRUSIVE ROCKS

—Most are associated with the Oligocene-Eocene Montezuma stock southwest of Silver Plume (Bookstrom and others, 1987). One dike appears to be related to Tertiary and Upper Cretaceous porphyries of the Empire stock northeast of Georgetown

- Td** Dacite (Oligocene and/or Eocene)—Bluish-gray to light-purple porphyry exhibiting conchoidal fracturing and generally characterized by altered feldspar phenocrysts in a groundmass of feldspar and quartz. Less commonly includes phenocrysts of quartz, hornblende, and biotite. Bookstrom (1993) suggested the dacite intrusives are 35 to 46 Ma

- Tgp** Granite porphyry (Oligocene and/or Eocene)—Tan to pinkish-gray, very fine-grained porphyry with numerous feldspar (orthoclase) phenocrysts and fewer biotite phenocrysts in a dense groundmass of quartz and orthoclase. Porphyry may be chalky yellow to orange where altered and is commonly extensively fractured. Occurs as small dikes and a larger stock northeast of Otter Mountain intruded 36.6 ± 4.2 Ma (Bookstrom and others, 1987)
- Tsp** Alaskite porphyry (Oligocene and/or Eocene)—White to light-gray porphyry composed of very fine-grained quartz and orthoclase with scattered orthoclase and plagioclase phenocrysts, and less commonly, quartz and mica. Dendrites of manganese oxide are pervasive along fracture and joint surfaces. Occurs as dikes typically less than 15 ft wide and as a large stock on the west flank of Paines Mountain intruded 37.0 ± 4 Ma (Bookstrom and others, 1987)

- Tqm** Quartz monzonite porphyry (Oligocene and/or Eocene)—Pale-violet to medium-gray porphyry composed of plagioclase, orthoclase, and quartz with phenocrysts of plagioclase and sandstone, and less commonly quartz and biotite. Occurs as dikes typically less than 15 ft wide and as a large stock on Lincoln Mountain. These intrusives are associated with the 35 to 40 Ma Montezuma stock (Bookstrom, 1993)

- Tkb** Bostonic porphyry (early Tertiary to Upper Cretaceous)—Single lilac to reddish-purple porphyry dike east of Sixon Mountain. Consists primarily of quartz and plagioclase with lesser amounts of quartz and amphibole. Rhombic orthoclase phenocrysts exhibit tan alteration halos extending outwards about 0.1 to 0.2 inches. This dike is similar in texture and mineralogy to rocks of the Empire stock, located northeast of the Georgetown quadrangle, which was emplaced between 61 to 68 Ma (Cunningham and others, 1994; Marvin and others, 1989; Simmons and Hodge, 1978)

PROTEROZOIC INTRUSIVE ROCKS—The majority of the Proterozoic intrusive rocks in the Georgetown quadrangle belong to the 1,400 Ma Berthoud plutonic suite (Tweto, 1987). Two small bodies of granodiorite at the northern end of the quadrangle belong to the 1,700 Ma Kourt plutonic suite (Tweto, 1987)

- Yd** Diorite and associated hornblende (Middle Proterozoic)—Dark-gray to mottled dark-gray and white, medium- to coarse-grained igneous rock composed primarily of plagioclase and hornblende with minor quartz and biotite. Associated greenish-black, fine- to medium-grained granular hornblende is defined almost entirely by hornblende. Spurr and others (1908) suggested the diorite and associated hornblende are slightly younger differentiation products of the magma that generated the granodiorite of the Mount Evans batholith

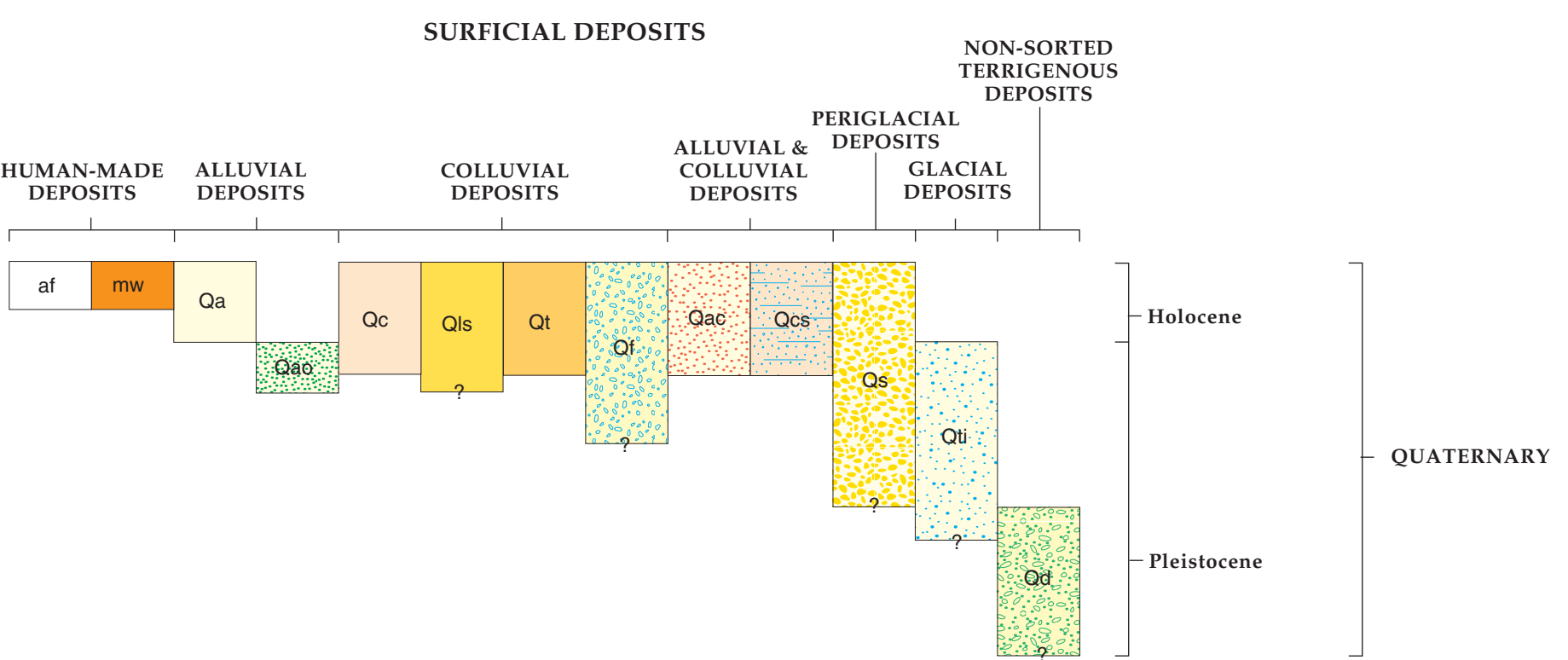
- Ysp** Silver Plume Granite (Middle Proterozoic)—Pink to pinkish-gray granite consisting primarily of microcline, plagioclase, and quartz, with minor to moderate amounts of biotite and/or muscovite. Abundant microcline phenocrysts are tabular and locally are weakly aligned. Where mineralized, off-white phenocrysts contrast with a rust- to purple-stained matrix. Age of the granite is 1,422 ± 2 Ma (Graubard and Mattison, 1990)

- Ygg** Biotite-muscovite granite and quartz monzonite (Middle Proterozoic)—Gray to pinkish-buff, fine- to medium-grained granite composed of microcline, plagioclase, and quartz, with lesser amounts of biotite and muscovite. May represent fine-grained equivalent of the Silver Plume Granite

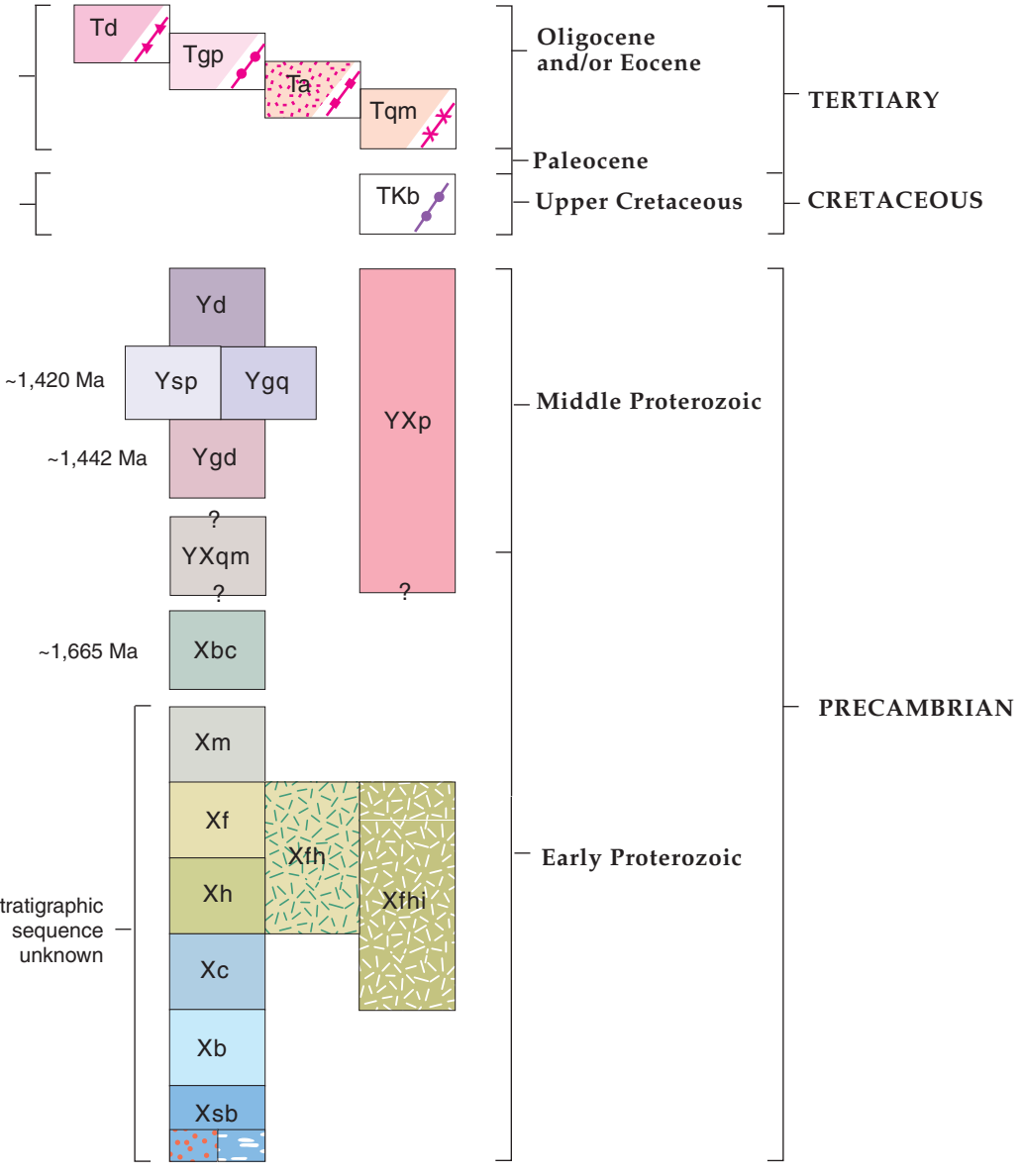
- Ygd** Granodiorite of the Mount Evans batholith (Middle Proterozoic)—Weakly to moderately foliated, medium- to coarse-grained, mottled black and white granodiorite composed of plagioclase, microcline, quartz, biotite, and accessory hornblende, magnetite, and sphene. Recent dating indicates the Mount Evans batholith was emplaced 1,422 ± 2 Ma and, therefore, is not equivalent to the Boulder Creek Granodiorite (Aleinikoff and others, 1993)

- YXgm** Quartz monzonite gneiss (Middle or Early Proterozoic)—Mottled black and white, medium- to coarse-grained, locally porphyritic gneiss composed essentially of plagioclase, orthoclase, and quartz with variable amounts of hornblende and biotite. Exhibits a gneissic structure defined by alternating bands of aligned quartz and feldspar and aligned biotite and hornblende. Locally intruded by dikes and small bodies of aplite and Silver Plume Granite

CORRELATION OF MAP UNITS



BEDROCK



- Yxp** Pegmatite, aplite, and related rocks (Middle and Early Proterozoic)—Includes pegmatite, granite pegmatite, and aplite, all of which consist of quartz and feldspar with various amounts of biotite, muscovite, magnetite, and hornblende. Granite pegmatite is medium-grained to pegmatitic. Aplite dikes and veins are composed of fine-grained quartz and feldspar

- Xbc** Boulder Creek Granodiorite (Early Proterozoic)—Medium- to coarse-grained, weakly foliated granodiorite exposed north of Empire Pass and composed of plagioclase, microcline, quartz, biotite, and accessory hornblende and sphene. Locally porphyritic with elongate or rounded, pinkish potassium feldspar crystals up to about 2 in. long. Age of the granodiorite is 1,665 ± 40 Ma (Peterman and others, 1968)

- Xm** Migmatite (Early Proterozoic)—Rocks that have been extensively intruded by granite or pegmatite, and/or intensely deformed and heated to the point of partial melting. Most commonly characterized by sillimanitic biotite gneiss (Xb) strongly intruded by Silver Plume Granite. Less commonly, includes layers or pods, biotite gneiss, hornblende gneiss, and felsic gneiss

- Xh** Hornblende gneiss and amphibolite (Early Proterozoic)—Medium- to dark-gray, fine-grained, massive or well-layered hornblende gneiss, and dark-green to black, fine- to medium-grained amphibolite. Hornblende gneiss is composed of hornblende and plagioclase and lesser amounts of quartz, biotite, and pyroxene. Amphibolite is composed almost entirely of hornblende and plagioclase with almost no quartz, biotite, or pyroxene

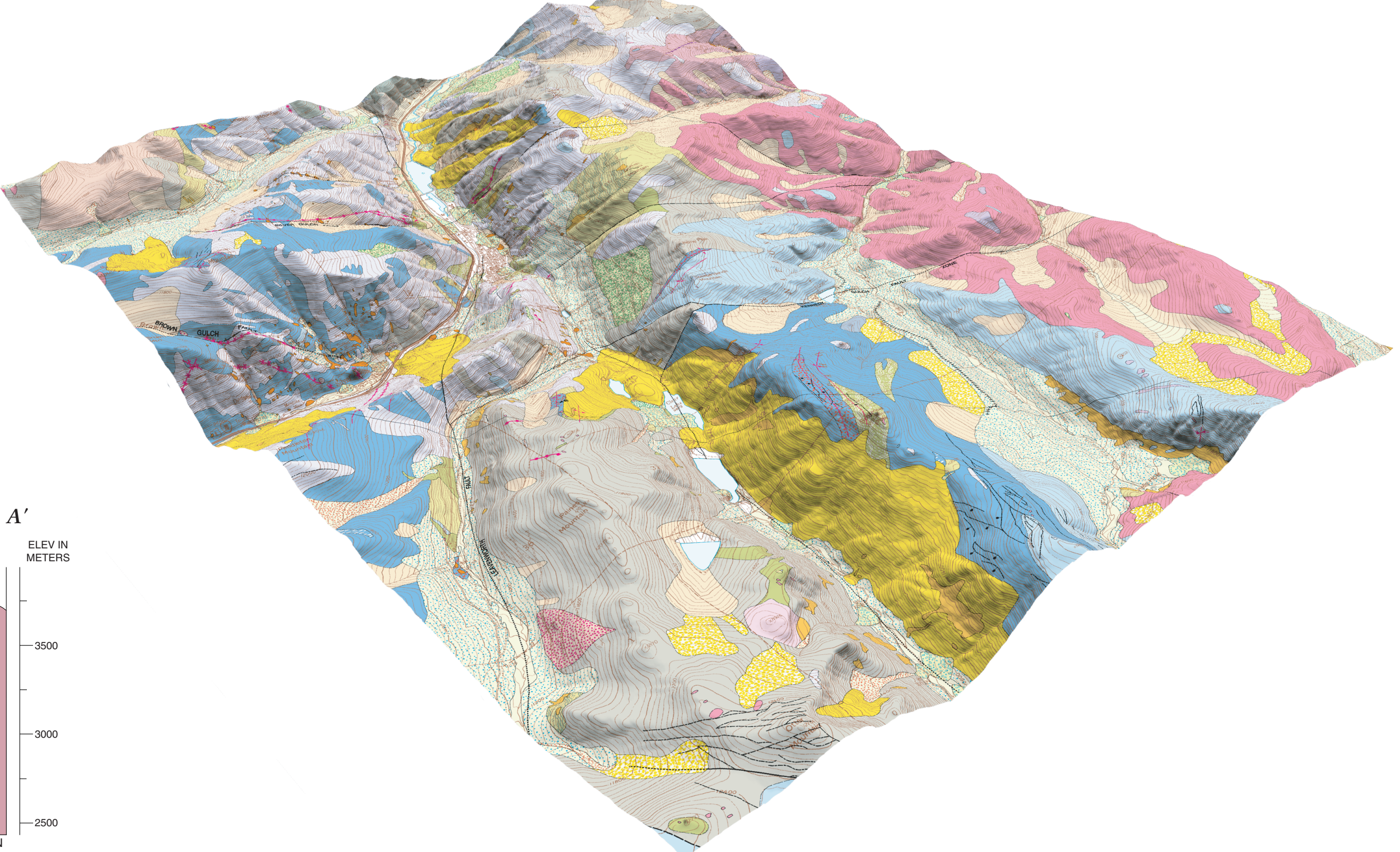
- Xc** Calc-silicate gneiss (Early Proterozoic)—Fine- to medium-grained, dark-gray to light-green gneiss consisting of quartz, feldspar, and hornblende, with accessory epidote, magnetite, and garnet. Commonly occurs as layers less than about 10 in. thick but locally may be several tens of feet thick

- Xh** Interlayered felsic and hornblende gneiss (Early Proterozoic)—Felsic gneiss (Xf) interlayered with hornblende gneiss (Xh). Thickness of individual layers is typically only a few feet but locally may exceed several tens of feet
- Xh** Interlayered felsic, hornblende, biotite, and calc-silicate gneiss (Early Proterozoic)—Felsic gneiss interlayered with hornblende gneiss, biotite gneiss, and sparse layers or pods of calc-silicate gneiss. Felsic and biotite gneiss layers range from a few inches to several tens of feet thick; hornblende and calc-silicate gneiss layers are usually less than about 18 in. thick

- Xb** Biotite gneiss (Early Proterozoic)—Fine-grained, light- to medium-gray gneiss composed primarily of biotite, quartz, and plagioclase, with minor magnetite, sillimanite, garnet, and/or cordierite. Typically equigranular and "salt and pepper" in appearance, though schistose in some areas

- Xsb** Sillimanitic biotite gneiss (Early Proterozoic)—Medium- to dark-gray, strongly foliated gneiss composed primarily of biotite, sillimanite, and quartz, with minor microcline, muscovite, garnet, tourmaline, and cordierite. Sillimanite is easily recognizable either as cloudy white rods and bundles or flattened pods elongate parallel to foliation. Open oval pattern indicates a high concentration of sillimanite pods generally greater than 1 in. and locally exceeding 8 in. in length. Pattern of red dots indicates a high concentration of macroscopic garnet crystals ranging from less than 0.1 in. up to about 1 in. in diameter

SHADED-RELIEF MAP OF THE GEORGETOWN QUADRANGLE WITH GEOLOGY AND TOPOGRAPHY OVERLAY, OBLIQUE VIEW LOOKING NORTHEAST



MAP SYMBOLS

- af** Geologic contact—Dashed where approximately located. Units shown as fractions where thin upper deposit overlies or occurs within lower unit
- af** Fault—Dashed where approximately located; dotted where concealed. Arrows show apparent direction of relative horizontal movement. Sense of movement on other faults is unknown
- af** Sacking—Linear scarp occurring on mountain ridges suggesting deep-seated rock creep, ridge-spreading movement, and slope failure in response to gradual gravitational movement of rock masses into adjacent valleys. Bar and ball on downthrown side of scarp
- af** Antiform—Showing axial trace and locally, plunge; dashed where approximately located; dotted where concealed
- af** Synform—Showing axial trace and locally, plunge; dashed where approximately located; dotted where concealed
- af** Trend and plunge of minor fold—Showing axial trace of minor folds in Proterozoic metamorphic rocks
- af** Moraine crest—Crest of prominent lateral and terminal moraines in Chicago Creek valley
- af** Strike and dip of foliation or compositional layering—Foliation is typically parallel to compositional layering in Proterozoic metamorphic rocks
- af** Inclined
- af** Vertical
- af** Strike and dip of joint
- af** Alignment of cross section



GEOLOGIC MAP OF THE GEORGETOWN QUADRANGLE, CLEAR CREEK COUNTY, COLORADO

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